

## 9. BASIC STRUCTURAL FEATURES

pointed out by Pandey & Krishna (1977) and suitable correction factors have recently been derived by Pandey, Prasad, Lele & Gauthier (1987). A satisfactory solution to the problem of structure determination of one-dimensionally disordered stackings must await proper understanding of all other factors that may influence the true diffraction profiles.

### 9.2.2. Layer stacking in general polytypic structures (By S. Durovič)

#### 9.2.2.1. The notion of polytypism

The common property of the structures described in Section 9.2.1 was the stacking ambiguity of adjacent layer-like structural units. This has been explained by the geometrical properties of close packing of equal spheres, and the different modifications thus obtained have been called *polytypes*.

This phenomenon was first recognized by Baumhauer (1912, 1915) as a result of his investigations of many SiC single crystals by optical goniometry. Among these, he discovered three *types* and his observations were formulated in five statements:

(1) all three types originate simultaneously in the same melt and seemingly also under the same, or nearly the same, conditions;

(2) they can be related in a simple way to the same axial ratio (each within an individual primary series);

(3) any two types (I and II, II and III) have certain faces in common but, except the basal face, there is no face occurring simultaneously in all three types;

(4) the crystals belonging to different, but also to all three, types often form intergrowths with parallel axes;

(5) any of the three types exhibits a typical X-ray diffraction pattern and thus also an individual molecular or atomic structure.

Baumhauer recognized the special role of these *types* within modifications of the same substance and called this phenomenon *polytypism* – a special case of polymorphism. The later determination of the crystal structures of Baumhauer's three types indicated that his results can be interpreted by a family of structures consisting of identical layers with hexagonal symmetry and differing only in their stacking mode.

The stipulation that the individual polytypes grow from the same system and under (nearly) the same conditions influenced for years the investigation of polytypes because it logically led to the question of their growth mechanism.

In the following years, many new polytypic substances have been found. Their crystal structures revealed that polytypism is restricted neither to close packings nor to heterodesmic 'layered structures' (e.g. CdI<sub>2</sub> or GaSe; cf. homodesmic SiC or ZnS; see §§9.2.1.2.2 to 9.2.1.2.4), and that the reasons for a stacking ambiguity lie in the crystal chemistry – in all cases the geometric nearest-neighbour relations between adjacent layers are preserved. The preservation of the bulk chemical composition was not questioned.

Some discomfort has arisen from refinements of the structures of various phyllosilicates. Here especially the micas exhibit a large variety of isomorphous replacements and it turns out that a certain chemical composition stabilizes certain polytypes, excludes others, and that the layers constituting polytypic structures need not be of the same kind. But subsequently the opinion prevailed that the sequence of individual kinds of layers in polytypes of the same family should remain the same and that the relative positions of adjacent layers cannot be completely random (e.g. Zvyagin, 1988). The postulates declared mixed-layer and turbostratic structures as non-polytypic. All this led to certain controversies about the notion of polytypism. While

Thompson (1981) regards polytypes as 'arising through different ways of stacking structurally compatible tabular units ... [provided that this] ... should not alter the chemistry of the crystal as a whole', Angel (1986) demands that 'polytypism arises from different modes of stacking of one or more structurally compatible modules', dropping thus any chemical constraints and allowing also for rod- and block-like modules.

The present official definition (Guinier *et al.*, 1984) reads:

“An element or compound is *polytypic* if it occurs in several different structural modifications, each of which may be regarded as built up by stacking layers of (nearly) identical structure and composition, and if the modifications differ only in their stacking sequence. Polytypism is a special case of polymorphism: the two-dimensional translations within the layers are (essentially) preserved whereas the lattice spacings normal to the layers vary between polytypes and are indicative of the stacking period. No such restrictions apply to polymorphism.

*Comment:* The above definition is designed to be sufficiently general to make polytypism a useful concept. There is increasing evidence that some polytypic structures are characterized either by small deviations from stoichiometry or by small amounts of impurities. (In the case of certain minerals like clays, micas and ferrites, deviations in composition up to 0.25 atoms per formula unit are permitted within the same polytypic series: two layer structures that differ by more than this amount should not be called polytypic.) Likewise, layers in different polytypic structures may exhibit slight structural differences and may not be isomorphic in the strict crystallographic sense.

The *Ad-Hoc* Committee is aware that the definition of polytypism above is probably too wide since it includes, for example, the turbostratic form of graphite as well as mixed-layer phyllosilicates. However, the sequence and stacking of layers in a polytype are always subject to well-defined limitations. On the other hand, a more general definition of polytypism that includes 'rod' and 'block' polytypes may become necessary in the future.”

This definition was elaborated as a compromise between members of the IUCr *Ad-Hoc* Committee on the Nomenclature of Disordered, Modulated and Polytype Structures. It is a slightly modified definition proposed by the IMA/IUCr Joint Committee on Nomenclature (Bailey *et al.*, 1977), which was the target of Angel's (1986) objections.

The official definition has indeed its shortcomings, but not so much in its restrictiveness concerning the chemical composition and structural rigidity of layers, because this can be overcome by a proper degree of abstraction (see below). More critical is the fact that it is not 'geometric' enough. It specifies neither the 'layers' (except for their two-dimensional periodicity), nor the limitations concerning their sequence and stacking mode, and it does not state the conditions under which a polytype belongs to a family.

Very impressive evidence that even polytypes that are in keeping with the first Baumhauer's statement may not have exactly the same composition and the structure of their constituting layers *cannot* be identical has been provided by studies on SiC carried out at the Leningrad Electrotechnical Institute (Sorokin, Tairov, Tsvetkov & Chernov, 1982; Tsvetkov, 1982). They indicate also that each periodic polytype is *sensu stricto* an individual polymorph. Therefore, it appears that the question whether some *real* polytypes belong to the same

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family depends mainly on the idealization and/or abstraction level, *relevant to a concrete purpose*.

This very idealization and/or abstraction process caused the term *polytype* to become also an *abstract notion* meaning a *structural type* with relevant geometrical properties,\* belonging to an abstract family whose members consist of layers with identical structure and keep identical bulk composition. Such an abstract notion lies at the root of all systemization and classification schemes of polytypes.

A still higher degree of abstraction has been achieved by Dornberger-Schiff (1964, 1966, 1979) who abstracted from chemical composition completely and investigated the manifestation of crystallochemical reasons for polytypism in the *symmetry of layers* and *symmetry relations between layers*. Her theory of OD (order-disorder) structures is thus a theory of symmetry of polytypes, playing here a role similar to that of group theory in traditional crystallography. In the next section, a brief account of basic terms, definitions, and logical constructions of OD theory will be given, together with its contribution to a geometrical definition of polytypism.

### 9.2.2.2. Symmetry aspects of polytypism

#### 9.2.2.2.1. Close packing of spheres

Polytypism of structures based on close packing of equal spheres (note this idealization) is explained by the fact that the spheres of any layer can be placed *either* in all the voids  $\nabla$  of the preceding layer, *or* in all the voids  $\Delta$  – not in both because of steric hindrance (Section 9.2.1, Fig. 9.2.1.1).

A closer look reveals that the two voids are *geometrically* (but not translationally) *equivalent*. This implies that the two possible pairs of adjacent layers, say *AB* and *AC*, are geometrically equivalent too – this equivalence is brought about *e.g.* by a reflection in any plane perpendicular to the layers and passing through the centres of mutually contacting spheres *A*: such a reflection transforms the layer *A* into itself, and *B* into *C*, and *vice versa*. Another important point is that the symmetry proper of any layer is described by the layer group  $P(6/m)mm$ ,<sup>†</sup> and that the relative position of any two adjacent layers is such that only some of the 24 symmetry operations of that layer group remain valid for the pair. It is easy to see that 12 out of the total of 24 transformations do not change the *z* coordinate of any starting point, and that these operations constitute a subgroup of the index [2]. These are the so-called  $\tau$  operations. The remaining 12 operations change any *z* into  $-z$ , thus turning the layer upside down; they constitute a coset. The latter are called  $\rho$  operations. Out of the 12  $\tau$  operations, only 6 are valid for the layer pair. One says that only these 6 operations have a *continuation* in the adjacent layer. Let us denote the general multiplicity of the group of  $\tau$  operations of a single layer by *N*, and that of the subgroup of these operations with a continuation in the adjacent layer by *F*: then the number *Z* of positions of the adjacent layer leading to geometrically equivalent layer pairs is given by  $Z = N/F$  (Dornberger-Schiff, 1964, pp. 32 ff.); in our case,  $Z = 12/6 = 2$  (Fig. 9.2.2.1). This is the so-called *NFZ relation*, valid with only minor alterations for all categories of OD structures (§9.2.2.2.7). It follows that all conceivable structures based on close packing of equal spheres are built on the same

symmetry principle: they consist of equivalent layers (*i.e.* layers of the same kind) and of equivalent layer pairs, and, in keeping with these stipulations, any layer can be stacked onto its predecessor in two ways. Keeping in mind that the layer pairs that are geometrically equivalent are also energetically equivalent, and neglecting in the first approximation the interactions between a given layer and the next-but-one layer, we infer that all structures built according to these principles are also energetically equivalent and thus equally likely to appear.

It is important to realize that the above symmetry considerations hold not only for close packing of spheres but also for *any conceivable structure* consisting of two-dimensionally periodic layers with symmetry  $P(6/m)mm$  and containing pairs of adjacent layers with symmetry  $P(3)m1$ . Moreover, the OD theory sets a quantitative stipulation for the relation between any two adjacent layers: they have to remain *geometrically equivalent* in any polytype belonging to a family. This is far more exact than the description: ‘the stacking of layers is such that it preserves the nearest-neighbour relationships’.

#### 9.2.2.2.2. Polytype families and OD groupoid families

All polytypes of a substance built on the same *structural principle* are said to belong to the same *family*. All polytypic structures, even of different substances, built according to the same *symmetry principle* also belong to a family, but different from the previous one since it includes structures of various polytype families, *e.g.* SiC, ZnS, AgI, which differ in their composition, lattice dimensions, *etc.* Such a family has been called an *OD groupoid family*; its members differ only in the relative distribution of *coincidence operations*\* describing the respective symmetries, irrespective of the crystallochemical content. These coincidence operations can be *total* or *partial* (local) and their set constitutes a groupoid (Dornberger-Schiff, 1964, pp. 16 ff.; Fichtner, 1965, 1977). Any polytype (abstract) belonging to such a family has its own stacking of layers, and its symmetry can be described by the appropriate individual groupoid. Strictly speaking, these groupoids are the members of an OD groupoid family. Let us recall that any space group

\*A coincidence operation is a space transformation (called also *isometric mapping*, *isometry*, or *motion*), which preserves distances between any two points of the given object.

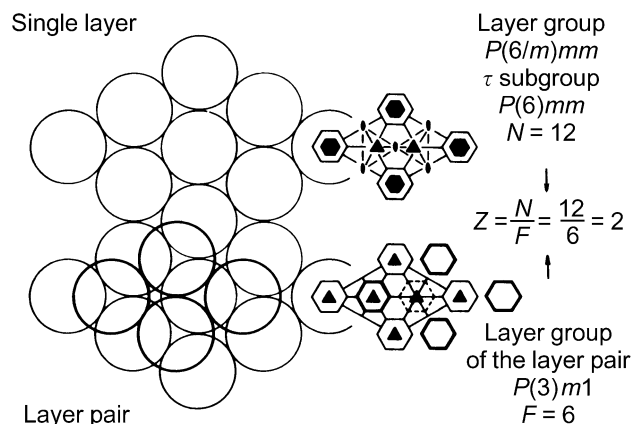


Fig. 9.2.2.1. Symmetry interpretation of close packings of equal spheres. The layer group of a single layer, the subgroup of its  $\tau$  operations, and the number of asymmetric units *N* per unit mesh of the former, are given at the top right. The  $\tau$  operations that have a continuation for the pair of adjacent layers, the layer group of the pair, and the value of *F* are indicated at the bottom right.

\*This is an interesting example of how a development in a scientific discipline influences semantics: *e.g.* when speaking of a  $6H$  polytype of SiC, one has very often in mind a characteristic sequence of Si-C layers rather than deviations from stoichiometry, presence and distribution of foreign atoms, distortion of coordination tetrahedra, *etc.*

<sup>†</sup>The direction in which there is no periodicity is indicated by parentheses (Dornberger-Schiff, 1959).

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consists of total coincidence operations only, which therefore become the symmetry operations for the entire structure.

### 9.2.2.2.3. MDO polytypes

Any family of polytypes theoretically contains an infinite number of periodic (Ross, Takeda & Wones, 1966; Mogami, Nomura, Miyamoto, Takeda & Sadanaga, 1978; McLarnan, 1981*a,b,c*) and non-periodic structures. The periodic polytypes, in turn, can again be subdivided into two groups, the 'privileged' polytypes and the remaining ones, and it depends on the approach as to how this is done. Experimentalists single out those polytypes that occur most frequently, and call them *basic*. Theorists try to predict basic polytypes, *e.g.* by means of geometrical and/or crystallochemical considerations. Such polytypes have been called *simple*, *standard*, or *regular*. Sometimes the agreement is very good, sometimes not. The OD theory pays special attention to those polytypes in which *all layer triples, quadruples, etc.*, are geometrically equivalent or, at least, which contain the smallest possible number of *kinds* of these units. They have been called polytypes with *maximum degree of order*, or *MDO polytypes*. The general philosophy behind the MDO polytypes is simple: all interatomic bonding forces decrease rapidly with increasing distance. Therefore, the forces between atoms of adjacent layers are decisive for the build-up of a polytype. Since the pairs of adjacent layers remain geometrically equivalent in all polytypes of a given family, these polytypes are in the first approximation also energetically equivalent. However, if the longer-range interactions are also considered, then it becomes evident that layer triples such as *ABA* and *ABC* in close-packed structures are, in general, energetically non-equivalent because they are also geometrically non-equivalent. Even though these forces are much weaker than those between adjacent layers, they may not be negligible and, therefore, under given crystallization conditions either one or the other kind of triples becomes energetically more favourable. It will occur again and again in the polytype thus formed, and not intermixed with the other kind. Such structures are – as a rule – sensitive to conditions of crystallization, and small fluctuations of these may reverse the energetical preferences, creating stacking faults and twinnings. This is why many polytypic substances exhibit non-periodicity.

As regards the close packing of spheres, the well known cubic and hexagonal polytypes *ABCABC...* and *ABAB...*, respectively, are MDO polytypes; the first contains only the triples *ABC*, the second only the triples *ABA*. Evidently, the MDO philosophy holds for a layer-by-layer rather than for a spiral growth mechanism. Since the symmetry principle of polytypic structures may differ considerably from that of close packing of equal spheres, the OD theory contains exact algorithms for the derivation of MDO polytypes in any category (Dornberger-Schiff, 1982; Dornberger-Schiff & Grell, 1982*a*).

### 9.2.2.2.4. Some geometrical properties of OD structures

As already pointed out, all relevant geometrical properties of a polytype family can be deduced from its symmetry principle. Let us thus consider a hypothetical simple family in which we shall disregard any concrete atomic arrangements and use geometrical figures with the appropriate symmetry instead.

Three periodic polytypes are shown in Fig. 9.2.2.2 (left-hand side). Any member of this family consists of equivalent layers perpendicular to the plane of the drawing, with symmetry  $P(1)m1$ . The symmetry of layers is indicated by isosceles triangles with a mirror plane  $[.m.]$ . All pairs of adjacent layers are also equivalent, no matter whether a layer is shifted by  $+b/4$

or  $-b/4$  relative to its predecessor, since the reflection across  $[.m.]$  transforms any given layer into itself and the adjacent layer from one possible position into the other. These two positions follow also from the *NFZ* relation:  $N = 2$ ,  $F = 1$  [the layer group of the pair of adjacent layers is  $P(1)11$ ] and thus  $Z = 2$ .

The layers are all equivalent and accordingly there must also be two coincidence operations transforming any layer into the adjacent one. The first operation is evidently the translation, the second is the glide reflection. If any of these becomes total for the remaining part of the structure, we obtain a polytype with all layer triples equivalent, *i.e.* a MDO polytype. The polytype (a) (Fig. 9.2.2.2) is one of them: the translation  $\mathbf{t} = \mathbf{a}_0 + \mathbf{b}/4$  is the total operation ( $|\mathbf{a}_0|$  is the distance between adjacent layers). It has basis vectors  $\mathbf{a}_1 = \mathbf{a}_0 + \mathbf{b}/4$ ,  $\mathbf{b}_1 = \mathbf{b}$ ,  $\mathbf{c}_1 = \mathbf{c}$ , space group  $P111$ , Ramsdell symbol  $1A$ ,\* Hägg symbol  $|+|$ . This polytype also has its enantiomorphous counterpart with Hägg symbol  $|-|$ . In the other polytype (b) (Fig. 9.2.2.2), the glide reflection is the total operation. The basis vectors of the polytype are

\* According to Guinier *et al.* (1984), triclinic polytypes should be designated  $A$  (anorthic) in their Ramsdell symbols.

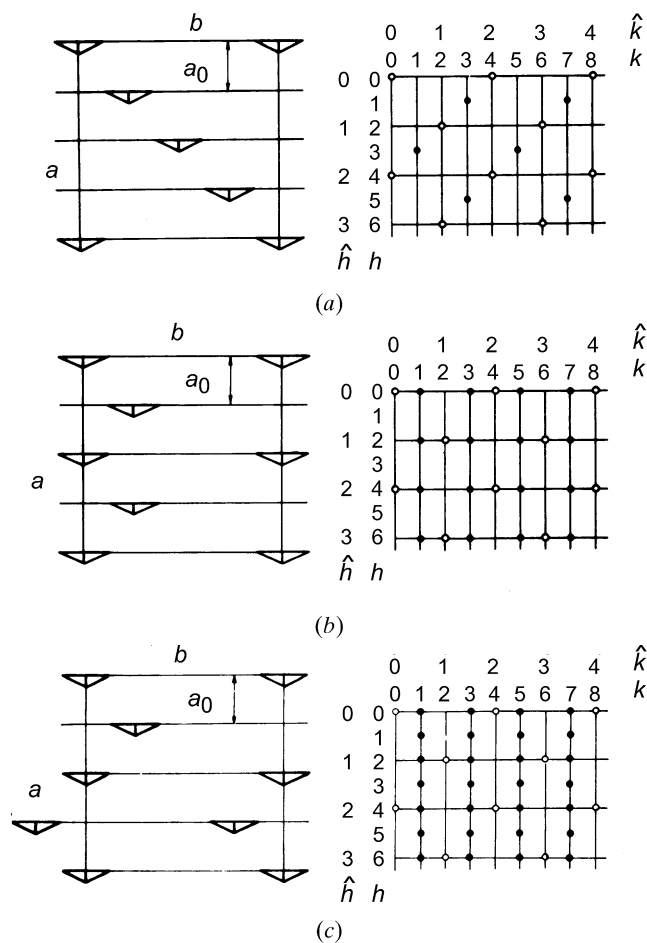


Fig. 9.2.2.2. Schematic representation of three structures belonging to the OD groupoid family  $P(1)m1|1$ ,  $y = 0.25$  (left). The layers are perpendicular to the plane of the drawing and their constituent atomic configurations are represented by isosceles triangles with symmetry  $[.m.]$ . All structures are related to a common orthogonal four-layer cell with  $\mathbf{a} = 4\mathbf{a}_0$ . The  $hk0$  nets in reciprocal space corresponding to these structures are shown on the right and the diffraction indices refer also to the common cell. Family diffractions common to all members of this family ( $k = 2\hat{k}$ ) and the characteristic diffractions for individual polytypes ( $k = 2\hat{k} + 1$ ) are indicated by open and solid circles, respectively.

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$\mathbf{a}_2 = 2\mathbf{a}_0$ ,  $\mathbf{b}_2 = \mathbf{b}$ ,  $\mathbf{c}_2 = \mathbf{c}$ , space group  $P1a1$ , Ramsdell symbol  $2M$ , Hagg symbol  $|+ -|$ . The equivalence of all layer triples in either of these polytypes is evident. The third polytype (c) (Fig. 9.2.2.2) is not a MDO polytype because it contains two kinds of layer triples, whereas it is possible to construct a polytype of this family containing only a selection of these. The polytype is again monoclinic with basis vectors  $\mathbf{a}_3 = 4\mathbf{a}_0$ ,  $\mathbf{b}_3 = \mathbf{b}$ ,  $\mathbf{c}_3 = \mathbf{c}$ , space group  $P1a1$ , Ramsdell symbol  $4M$ , and Hagg symbol  $|+ - - +|$ .

Evidently, the partial mirror plane is crucial for the polytypism of this family. And yet the space group of none of its periodic members can contain it – simply because it can never become total. The space-group symbols thus leave some of the most important properties of periodic polytypes unnoticed. Moreover, the atomic coordinates of different polytypes expressed in terms of the respective lattice geometries cannot be immediately compared. And, finally, for non-periodic members of a family, a space-group symbol cannot be written at all. This is why the OD theory gives a *special symbol* indicating the symmetry proper of individual layers ( $\lambda$  symmetry) as well as the coincidence operations transforming a layer into the adjacent one ( $\sigma$  symmetry). The symbol of the OD groupoid family of our hypothetical example thus consists of two lines (Dornberger-Schiff, 1964, pp. 41 ff.; Fichtner, 1979a,b):

$$\begin{array}{llll} P(1) & m & 1 & \lambda \text{ symmetry} \\ \{(1) & a_2 & 1\} & \sigma \text{ symmetry,} \end{array}$$

where the unusual subscript 2 indicates that the glide reflection transforms the given layer into the subsequent one.

It is possible to write such a symbol for any OD groupoid family for equivalent layers, and thus also for the close packing of spheres. However, keeping in mind that the number of asymmetric units here is 24 ( $\lambda$  symmetry), one has to indicate also 24  $\sigma$  operations, which is instructive but unwieldy. This is why Fichtner (1980) proposed simplified one-line symbols, containing full  $\lambda$  symmetry and only the rotational part of *any one* of the  $\sigma$  operations plus its translational components. Accordingly, the symbol of our hypothetical family reads:  $P(1)m1|1$ ,  $y = 0.25$ ; for the family of close packings of equal spheres:  $P(6/m)mm|1$ ,  $x = 2/3$ ,  $y = 1/3$  (the layers are in both cases translationally equivalent and the rotational part of a translation is the identity).

An OD groupoid family symbol should not be confused with a *polytype symbol*, which gives information about the structure of an individual polytype (Dornberger-Schiff, urovic & Zvyagin, 1982; Guinier *et al.*, 1984).

### 9.2.2.2.5. Diffraction pattern – structure analysis

Let us now consider schematic diffraction patterns of the three structures on the right-hand side of Fig. 9.2.2.2. It can be seen that, while being in general different, they contain a common subset of diffractions with  $k = 2\hat{k}$  – these, normalized to a constant number of layers, have the same distribution of intensities and monoclinic symmetry. This follows from the fact that they correspond to the so-called *superposition structure* with basis vectors  $\mathbf{A} = 2\mathbf{a}_0$ ,  $\mathbf{B} = \mathbf{b}/2$ ,  $\mathbf{C} = \mathbf{c}$ , and space group  $C1m1$ . It is a fictitious structure that can be obtained from any of the structures in Fig. 9.2.2.2 as a normalized sum of the structure in its given position and in a position shifted by  $b/2$ , thus

$$\hat{\rho}(xyz) = \frac{1}{2}[\rho(xyz) + \rho(x, y + 1/2, z)].$$

Evidently, this holds for all members of the family, including the non-periodic ones. In general, the superposition structure is obtained by simultaneous realization of all  $Z$  possible positions of all OD layers in any member of the family (Dornberger-Schiff, 1964, p. 54). As a consequence, its symmetry can be obtained by completing any of the family groupoids to a group (Fichtner, 1977). This structure is by definition periodic and *common to all members* of the family. Thus, the corresponding diffractions are also always sharp, common, and characteristic for the family. They are called *family diffractions*.

Diffractions with  $k = 2\hat{k} + 1$  are characteristic for individual members of the family. They are sharp for periodic polytypes but appear as diffuse streaks for non-periodic ones. Owing to the  $C$  centring of the superposition structure, only diffractions with  $\hat{h} + \hat{k} = 2n$  are present. It follows that  $0kl$  diffractions are present only for  $\hat{k} = 2n$ , which, in an indexing referring to the actual  $\mathbf{b}$  vector reads:  $0kl$  present only for  $k = 4n$ . This is an example of non-space-group absences exhibited by many polytypic structures. They can be used for the determination of the OD groupoid family (Dornberger-Schiff & Fichtner, 1972).

There is no routine method for the determination of the structural principle of an OD structure. It is easiest when one has at one's disposal many different (at least two) periodic polytypes of the same family with structures solved by current methods. It is then possible to compare these structures, determine equivalent regions in them (Grell, 1984), and analyse partial symmetries. This results in an OD interpretation of the substance and a description of its polytypism.

Sometimes it is possible to arrive at an OD interpretation from one periodic structure, but this necessitates experience in the recognition of the partial symmetry and prediction of potential polytypism (Merlino, Orlandi, Perchiazzi, Basso & Palenzona, 1989).

The determination of the structural principle is complex if only disordered polytypes occur. Then – as a rule – the superposition structure is solved first by current methods. The actual structure of layers and relations between them can then be determined from the intensity distribution along diffuse streaks (for more details and references see Jagodzinski, 1964; Sedlacek, Kuban & Backhaus, 1987; Muller & Conradi, 1986). High-resolution electron microscopy can also be successfully applied – see Subsection 9.2.2.4.

### 9.2.2.2.6. The vicinity condition

A polytype family contains periodic as well as non-periodic members. The latter are as important as the former, since the very fact that they can be non-periodic carries important crystallochemical information. Non-periodic polytypes do not comply with the classical definition of crystals, but we believe that this definition should be generalized to include rather than exclude non-periodic polytypes from the world of crystals (Dornberger-Schiff & Grell, 1982b). The OD theory places them, together with the periodic ones, in the hierarchy of the so-called *VC structures*. The reason for this is that all periodic structures, even the non-polytypic ones, can be thought of as consisting of *disjunct*, two-dimensionally periodic slabs, the *VC layers*, which are stacked together according to three rules called the *vicinity condition* (VC) (Dornberger-Schiff, 1964, pp. 29 ff., 1979; Dornberger-Schiff & Fichtner, 1972):

( $\alpha$ ) VC layers are either geometrically equivalent or, if not, they are relatively few in kind;

( $\beta$ ) translation groups of all VC layers are either identical or they have a common subgroup;

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( $\gamma$ ) equivalent sides of equivalent layers are faced by equivalent sides of adjacent layers so that the resulting pairs are equivalent [for a more detailed specification and explanation see Dornberger-Schiff (1979)].

If the stacking of VC layers is unambiguous, traditional three-dimensionally periodic structures result (*fully ordered structures*). *OD structures* are VC structures in which the stacking of VC layers is ambiguous at every layer boundary ( $Z > 1$ ). The corresponding VC layers then become *OD layers*. OD layers are, in general, not identical with crystallochemical layers; they may contain half-atoms at their boundaries. In this context, they are analogous with unit cells in traditional crystallography, which may also contain parts of atoms at their boundaries. However, *the choice of OD layers is not absolute*: it depends on the polytypism, either actually observed or reasonably anticipated, on the degree of symmetry idealization, and other circumstances (Grell, 1984).

### 9.2.2.2.7. Categories of OD structures

Any OD layer is two-dimensionally periodic. Thus, a unit mesh can be chosen according to the conventional rules for the corresponding layer group; the corresponding vectors or their linear combinations (Zvyagin & Fichtner, 1986) yield the basis vectors parallel to the layer plane and thus also their lengths as units for fractional atomic coordinates. But, in general, there is no periodicity in the direction perpendicular to the layer plane and it is thus necessary to define the corresponding unit length in some other way. This depends on the symmetry principle of the family in question – or, more narrowly, on the *category* to which this family belongs.

OD structures can be built of *equivalent layers* or contain *layers of several kinds*. The rule ( $\gamma$ ) of the VC implies that a projection of any OD structure – periodic or not – on the stacking direction is periodic. This period, called *repeat unit*, is the required unit length.

#### 9.2.2.2.7.1. OD structures of equivalent layers

If the OD layers are equivalent then they are either all polar or all non-polar in the stacking direction. Any two adjacent polar layers can be related either by  $\tau$  operations only, or by  $\rho$  operations only. For non-polar layers, the  $\sigma$  operations are both  $\tau$  and  $\rho$ . Accordingly, there are *three categories* of OD structures of equivalent layers. They are shown schematically in Fig. 9.2.2.3; the character of the corresponding  $\lambda$  and  $\sigma$  operations is as follows (Dornberger-Schiff, 1964, pp. 24 ff.):

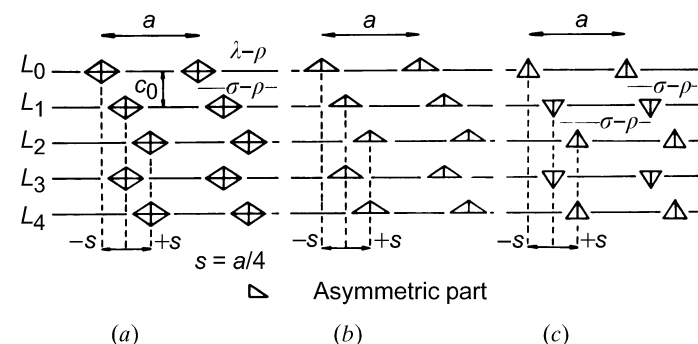


Fig. 9.2.2.3. Schematic examples of the three categories of OD structures consisting of equivalent layers (perpendicular to the plane of the drawing): (a) category I – OD layers non-polar in the stacking direction; (b) category II – polar OD layers, all with the same sense of polarity; (c) category III – polar OD layers with regularly alternating sense of polarity. The position of  $\rho$  planes is indicated.

	category I	category II	category III
$\lambda$ operations	$\tau$ and $\rho$	$\tau$	$\tau$
$\sigma$ operations	$\tau$ and $\rho$	$\tau$	$\rho$

Category II is the simplest: the OD layers are polar and all with the same sense of polarity (they are  $\tau$ -equivalent); our hypothetical example given in §9.2.2.2.4 belongs to this category. The layers can thus exhibit only one of the 17 polar layer groups. The projection of any vector between two  $\tau$ -equivalent points in two adjacent layers on the stacking direction (perpendicular to the layer planes) is the repeat unit and it is denoted by  $\mathbf{c}_0$ ,  $\mathbf{a}_0$ , or  $\mathbf{b}_0$  depending on whether the basis vectors in the layer plane are  $\mathbf{ab}$ ,  $\mathbf{bc}$ , or  $\mathbf{ca}$ , respectively. The choice of origin in the stacking direction is arbitrary but preferably so that the  $z$  coordinates of atoms within a layer are positive. Examples are SiC, ZnS, and AgI.

OD layers in category I are non-polar and they can thus exhibit any of the 63 non-polar layer groups. Inspection of Fig. 9.2.2.3(a) reveals that the symmetry elements representing the  $\lambda$ - $\rho$  operations (*i.e.* the operations turning a layer upside down) can lie only in one plane called the *layer plane*. Similarly, the symmetry elements representing the  $\sigma$ - $\rho$  operations (*i.e.* the operations converting a layer into the adjacent one) also lie in one plane, located exactly halfway between two nearest layer planes. These two kinds of planes are called  $\rho$  planes. The distance between two nearest layer planes is the repeat unit  $c_0$ . Examples are close packing of equal spheres, GaSe,  $\alpha$ -wollastonite (Yamanaka & Mori, 1981),  $\beta$ -wollastonite (Ito, Sadanaga, Takéuchi & Tokonami, 1969),  $\text{K}_3[\text{M}(\text{CN})_6]$  (Jagner, 1985), and many others.

The OD structures belonging to the above two categories contain pairs of adjacent layers, all equivalent. This does not apply for structures of category III, which consist of polar layers that are converted into their neighbours by  $\rho$  operations. It is evident (Fig. 9.2.2.3c) that two kinds of pairs of adjacent layers are needed to build any such structure. It follows that only even-numbered layers can be mutually  $\tau$ -equivalent and the same holds for odd-numbered layers. There are only  $\sigma$ - $\rho$  planes in these structures, and again they

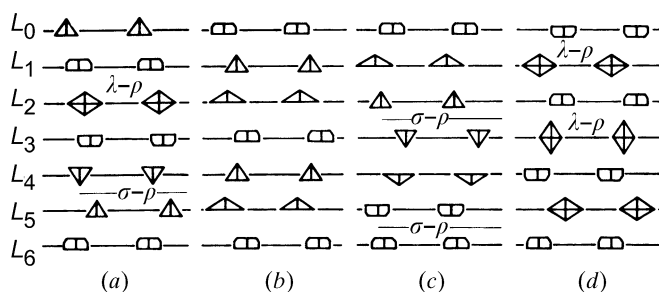


Fig. 9.2.2.4. Schematic examples of the four categories of OD structures consisting of more than one kind of layer (perpendicular to the plane of the drawing). Equivalent OD layers are represented by equivalent symbolic figures. (a) Category I – three kinds of OD layers: one kind ( $L_{2+5n}$ ) is non-polar, the remaining two are polar. One and only one kind of non-polar layer is possible in this category. (b) Category II – three kinds of polar OD layers; their triples are polar and retain their sense of polarity in the stacking direction. (c) Category III – three kinds of polar OD layers; their triples are polar and regularly change their sense of polarity in the stacking direction. (d) Category IV – three kinds of OD layers: two kinds are non-polar ( $L_{1+4n}$  and  $L_{3+4n}$ ), one kind is polar. Two and only two kinds of non-polar layers are possible in this category. The position of  $\rho$  planes is indicated.

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are of two kinds; the origin can be placed in either of them.  $c_0$  is the distance between two nearest  $\rho$  planes of the same kind, and slabs of this thickness contain two OD layers. There are three examples for this category known to date: foshagite (Gard & Taylor, 1960),  $\gamma$ -Hg<sub>3</sub>S<sub>2</sub>Cl<sub>2</sub> (Đurovič, 1968), and 2,2-aziridinedicarboxamide (Fichtner & Grell, 1984).

### 9.2.2.2.7.2. OD structures with more than one kind of layer

If an OD structure consists of  $N > 1$  kinds of OD layers, then it can be shown (Dornberger-Schiff, 1964, pp. 64 ff.) that it can fall into one of *four categories*, according to the polarity or non-polarity of its constituent layers and their sequence. These are shown schematically in Fig. 9.2.2.4; the character of the corresponding  $\lambda$  and  $\sigma$  operations is

	category I	category II	category III	category IV
$\lambda$ operations	$\tau$ and $\rho$ (one set) $\tau$ ( $N - 1$ sets)	$\tau$ ( $N$ sets)	$\tau$ ( $N$ sets)	$\tau$ and $\rho$ (two sets) $\tau$ ( $N - 2$ sets)
$\sigma$ operations	$\rho$ (one set)	none	$\rho$ (two sets)	none.

Here also category II is the simplest. The structures consist of  $N$  kinds of cyclically recurring polar layers whose sense of polarity remains unchanged (Fig. 9.2.2.4b). The choice of origin in the stacking direction is arbitrary;  $c_0$  is the projection on this direction of the shortest vector between two  $\tau$ -equivalent points – a slab of this thickness contains all  $N$  OD layers of different kinds. Examples are the structures of the serpentine-kaolin group.

Structures of category III also consist of polar layers but, in contrast to category II, the  $N$ -tuples containing all  $N$  different OD layers each alternate regularly the sense of their polarity in the stacking direction. Accordingly (Fig. 9.2.2.4c), there are two kinds of  $\sigma$ - $\rho$  planes and two kinds of pairs of equivalent adjacent layers in these structures. The origin can be placed in either of the two  $\rho$  planes.  $c_0$  is the distance between the nearest two equivalent  $\rho$  planes; a slab with this thickness contains  $2 \times N$  non-equivalent OD layers. No representative of this category is known to date.

The structures of category I contain one, and only one, kind of non-polar layer, the remaining  $N - 1$  kinds are polar and alternate in their sense of polarity along the stacking direction (Fig. 9.2.2.4a). Again, there are two kinds of  $\rho$  planes here, but one is a  $\lambda$ - $\rho$  plane (the layer plane of the non-polar OD layer), the other is a  $\sigma$ - $\rho$  plane. These structures thus contain only one kind of pair of equivalent adjacent layers. The origin is placed in the  $\lambda$ - $\rho$  plane.  $c_0$  is the distance between the nearest two equivalent  $\rho$  planes and a slab with this thickness contains  $2 \times (N - 1)$  non-equivalent polar OD layers plus one entire non-polar layer. Examples are the  $MX_2$  compounds (CdI<sub>2</sub>, MoS<sub>2</sub>, etc.) and the talc-pyrophyllite group.

The structures of category IV contain two, and only two, kinds of non-polar layers. The remaining  $N - 2$  kinds are polar and alternate in their sense of polarity along the stacking direction (Fig. 9.2.2.4d). Both kinds of  $\rho$  planes are  $\lambda$ - $\rho$  planes, identical with the layer planes of the non-polar OD layers; the origin can be placed in any one of them.  $c_0$  is chosen as in categories I and III. A slab with this thickness contains  $2 \times (N - 2)$  non-equivalent polar layers plus the two non-polar layers. Examples are micas, chlorites, vermiculites, etc.

OD structures containing  $N > 1$  kinds of layers need special symbols for their OD groupoid families (Grell & Dornberger-Schiff, 1982).

A slab of thickness  $c_0$  containing the  $N$  non-equivalent polar OD layers in the sequence as they appear in a given structure of category II represents completely its composition. In the remaining three categories, a slab with thickness  $c_0/2$ , the polar part of the structure contained between two adjacent  $\rho$  planes, suffices. Such slabs are higher structural units for OD structures of more than one kind of layer and have been called *OD packets*. An OD packet is thus defined as the smallest continuous part of an OD structure that is periodic in two dimensions and which represents its composition completely (Đurovič, 1974a).

The hierarchy of VC structures is shown in Fig. 9.2.2.5.

### 9.2.2.2.8. Desymmetrization of OD structures

If a fully ordered structure is refined, using the space group determined from the systematic absences in its diffraction pattern and then by using some of its subgroups, serious discrepancies are only rarely encountered. Space groups thus characterize the general symmetry pattern quite well, even in real crystals. However, experience with refined periodic polytypic structures has revealed that there are always significant deviations from the OD symmetry and, moreover, even the atomic coordinates within OD layers in different polytypes of the same family may differ from one another. The OD symmetry thus appears as only an approximation to the actual symmetry pattern of polytypes. This phenomenon was called *desymmetrization* of OD structures (Đurovič, 1974b, 1979).

When trying to understand this phenomenon, let us recall the structure of rock salt. Its symmetry  $Fm\bar{3}m$  is an expression of the energetically most favourable relative position of Na<sup>+</sup> and Cl<sup>-</sup> ions in this structure – the right angles  $\alpha\beta\gamma$  follow from the symmetry. Since the whole structure is cubic, we cannot expect that the environment of any building unit, e.g. of any octahedron NaCl<sub>6</sub>, would exercise on it an influence that would decrease its symmetry; the symmetries of these units and of the whole structure are not ‘antagonistic’.

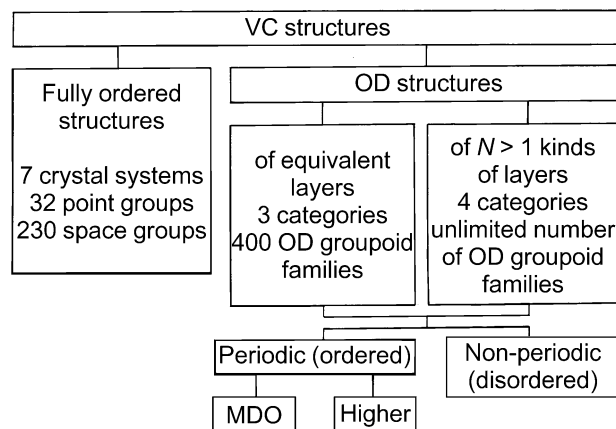


Fig. 9.2.2.5. Hierarchy of VC structures indicating the position of OD structures within it.

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Not so in OD structures, where any OD layer is by definition situated in a disturbing environment because its symmetry does not conform to that of the entire structure. 'Antagonistic' relations between these symmetries are most drastic in pure MDO structures because of the regular sequence of layers. The partial symmetry operations become irrelevant and the OD groupoid degenerates into the corresponding space group.

The more disordered an OD structure is, the smaller become the disturbing effects that the environment exercises on an OD layer. These can be, at least statistically, neutralized by random positions of neighbouring layers so that partial symmetry operations can retain their relevance throughout the structure. This can be expressed in the form of a paradox: the less periodic an OD structure is, the more symmetric it appears.

Despite desymmetrization, the OD theory remains a geometrical theory that can handle properly the *general symmetry pattern* of polytypes (which group theory cannot). It establishes a *symmetry norm* with which deviations observed in real polytypes can be compared. Owing to the high abstraction power of OD considerations, systematics of entire families of polytypes at various degree-of-idealization levels can be worked out, yielding thus a common point of view for their treatment.

### 9.2.2.2.9. Concluding remarks

Although very general physical principles (OD philosophy, MDO philosophy) underlie the OD theory, it is mainly a geometrical theory, suitable for a *description* of the symmetry of polytypes and their families rather than for an *explanation* of polytypism. It thus does not compete with crystal chemistry, but cooperates with it, in analogy with traditional crystallography, where group theory does not compete with crystal chemistry.

When speaking of polytypes, one should always be aware, whether one has in mind a concrete *real* polytype – more or less in Baumhauer's sense – or an *abstract* polytype as a structural type (Subsection 9.2.2.1).

A substance can, in general, exist in the form of various polymorphs and/or polytypes of one or several families. Since polytypes of the same family differ only slightly in their crystal energy (Verma & Krishna, 1966), an entire family can be considered as an energetic analogue to one polymorph. As a rule, polytypes belonging to different families of the same substance do not co-exist.  $\text{Al}(\text{OH})_3$  may serve as an example for two different families: the bayerite family, in which the adjacent planes of OH groups are stacked according to the principle of close packing (Zvyagin *et al.*, 1979), and the gibbsite-nordstrandite family in which these groups coincide in the normal projection.\* Another example is the phyllosilicates (§9.2.2.3.1). The compound  $\text{Hg}_3\text{S}_2\text{Cl}_2$ , on the other hand, is known to yield two polymorphs  $\alpha$  and  $\beta$  (Carlson, 1967; Frueh & Gray, 1968) and one OD family of  $\gamma$  structures (Đurovič, 1968).

As far as the definition of *layer polytypism* is concerned, OD theory can contribute specifications about the layers themselves and the geometrical rules for their stacking within a family (all incorporated in the vicinity condition). A possible definition might then read:

*Polytypism* is a special case of polymorphism, such that the individual polymorphs (called *polytypes*) may be regarded as arising through different modes of stacking layer-like structural

\* Sandwiches with composition  $\text{Al}(\text{OH})_3$  (similar to those in  $\text{CdI}_2$ ) are the same in both families, but their stacking mode is different. This and similar situations in other substances might have been the reason for distinguishing between 'polytype diversity' and 'OD diversity' (Zvyagin, 1988).

units. The layers and their stackings are limited by the *vicinity condition*. All polytypes built on the same structural principle belong to a *family*; this depends on the degree of a structural and/or compositional *idealization*.

Geometrical theories concerning rod and block polytypism have not yet been elaborated, the main reason is the difficulty of formulating properly the vicinity condition (Sedlacek, Grell & Dornberger-Schiff, private communications). But such structures are known. Examples are the structures of tobermorite (Hamid, 1981) and of manganese(III) hydrogenbis(orthophosphate) dihydrate (Čísařová, Novák & Petříček, 1982). Both structures can be thought of as consisting of a three-dimensionally periodic framework of certain atoms into which one-dimensionally periodic chains and aperiodic finite configurations of the remaining atoms, respectively, 'fit' in two equivalent ways.

### 9.2.2.3. Examples of some polytypic structures

The three examples below illustrate the three main methods of analysis of polytypism indicated in §9.2.2.2.5.

#### 9.2.2.3.1. Hydrous phyllosilicates

The basic concepts were introduced by Pauling (1930*a,b*) and confirmed later by the determination of concrete crystal structures. A *crystallochemical analysis* of these became the basis for generalizations and systemizations. The aim was the understanding of geometrical reasons for the polytypism of these substances as well as the development of identification routines through the derivation of basic polytypes (§9.2.2.2.3). Smith & Yoder (1956) succeeded first in deriving the six basic polytypes in the mica family.

Since the 1950's, two main schools have developed: in the USA, represented mainly by Brindley, Bailey, and their co-workers (for details and references see Bailey, 1980, 1988*a*; Brindley, 1980), and in the former USSR, represented by Zvyagin and his co-workers (for details and references see Zvyagin, 1964, 1967; Zvyagin *et al.*, 1979). Both these schools based their systemizations on idealized structural models corresponding to the ideas of Pauling, with *hexagonal* symmetry of tetrahedral sheets (see later). The US school uses indicative symbols (Guinier *et al.*, 1984) for the designation of individual polytypes, and single-crystal as well as powder X-ray diffraction methods for their identification, whereas the USSR school uses unitary descriptive symbols for polytypes of all mineral groups and mainly electron diffraction on oblique textures for identification purposes. For the derivation of basic polytypes, both schools use crystallochemical considerations; symmetry principles are applied tacitly rather than explicitly.

In contrast to crystal structures based on close packings, where all relevant details of individual (even multilayer) polytypes can be recognized in the (1120) section, the structures of hydrous phyllosilicates are rather complex. For their representation, Figueiredo (1979) used the concept of *condensed models*.

Since 1970, the OD school has also made its contribution. In a series of articles, basic types of hydrous phyllosilicates have been interpreted as OD structures of  $N > 1$  kinds of layers: the serpentine-kaolin group (Dornberger-Schiff & Đurovič, 1975*a,b*), Mg-vermiculite (Weiss & Đurovič, 1980), the mica group (Dornberger-Schiff, Backhaus & Đurovič, 1982; Backhaus & Đurovič, 1984; Đurovič, Weiss & Backhaus, 1984; Weiss & Wiewióra, 1986), the talc-pyrophyllite group (Đurovič & Weiss, 1983; Weiss & Đurovič, 1985*a*), and the chlorite group (Đurovič, Dornberger-Schiff & Weiss, 1983;

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Weiss & Āuroviĉ, 1983). The papers published before 1983 use the Pauling model; the later papers are based on the model of Radoslovich (1961) with *trigonal* symmetry of tetrahedral sheets. In all cases, MDO polytypes (§9.2.2.3) have been derived systematically: their sets partially overlap with basic polytypes presented by the US or the USSR schools. The OD models allowed the use of unitary descriptive symbols for individual polytypes from which all the relevant symmetries can be determined (Āuroviĉ & Dornberger-Schiff, 1981) as well as of extended indicative Ramsdell symbols (Weiss & Āuroviĉ, 1985b). The results, including principles for identification of polytypes, have been summarized by Āuroviĉ (1981).

The main features of polytypes of basic types of hydrous phyllosilicates, of their diffraction patterns and principles for their identification, are given in the following.

### 9.2.2.3.1.1. General geometry

Tetrahedral and octahedral sheets are the fundamental, *two-dimensionally periodic* structural units, common to all hydrous phyllosilicates. Any *tetrahedral sheet* consists of  $(\text{Si,Al,Fe}^{3+},\text{Ti}^{4+})\text{O}_4$  tetrahedra joined by their three basal O atoms to form a network with symmetry  $P(3)1m$  (Fig. 9.2.2.6a). The atomic coordinates can be related either to a hexagonal axial system with a primitive unit mesh and basis vectors  $\mathbf{a}_1, \mathbf{a}_2$ , or to an orthohexagonal system with a *c*-centred unit mesh and basis vectors  $\mathbf{a}, \mathbf{b}$  ( $b = \sqrt{3}a$ ). Any *octahedral sheet* consists of  $M(\text{O,OH})_6$  octahedra with shared edges (Fig. 9.2.2.6b), and with cations *M* most frequently  $\text{Mg}^{2+}, \text{Al}^{3+}, \text{Fe}^{2+}, \text{Fe}^{3+}$ , but also  $\text{Li}^+, \text{Mn}^{2+}$  ( $r_M < \sim 0.9 \text{ \AA}$ ), etc. There are three octahedral

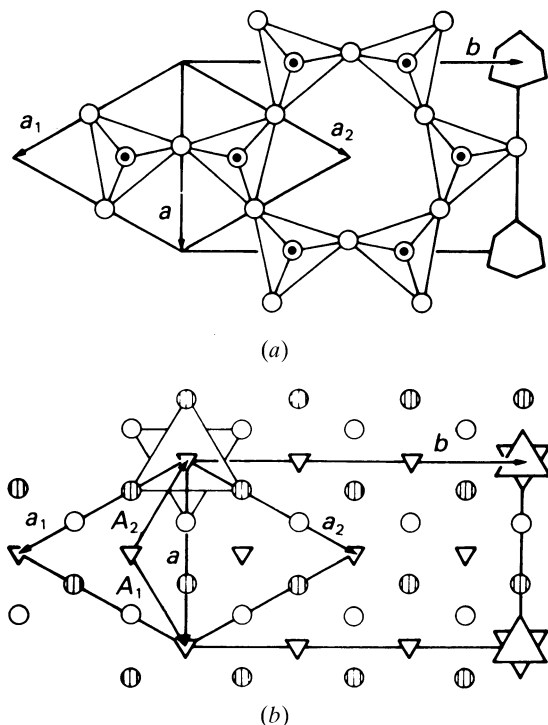


Fig. 9.2.2.6. (a) Tetrahedral sheet in a normal projection. Open circles: basal oxygen atoms, circles with black dots: apical oxygen atoms and tetrahedral cations. Hexagonal and orthohexagonal basis vectors and symbolic figures (ditrigons) for pictorial representation of these sheets are also shown. (b) Octahedral sheet. Open and shaded circles belong to the lower and the upper oxygen atomic planes, respectively; small triangles denote octahedral sites. Triangular stars on the right are the symbolic figures for pictorial representation of these sheets: the two triangles correspond to the lower and upper basis of any octahedron, respectively.

sites per unit mesh  $\mathbf{a}_1, \mathbf{a}_2$ . Crystallochemical classification distinguishes between two extreme cases: *trioctahedral* (all three octahedral sites are occupied) and *dioctahedral* (one site is – even statistically – empty). This classification is based on a bulk chemical composition. A classification from the symmetry point of view distinguishes between three cases: *homo-octahedral* [all three octahedral sites are occupied by the same kind of crystallochemical entity, *i.e.* either by the same kind of ion or by a statistical average of different kinds of ions including voids; symmetry of such a sheet is  $H(3)12/m$ ]\*; *meso-octahedral* [two octahedral sites are occupied by the same kind of crystallochemical entity, the third by a different one, in an ordered way; symmetry  $P(3)12/m$ ]; and *hetero-octahedral* [each octahedral site is occupied by a different crystallochemical entity in an ordered way; symmetry  $P(3)12$ ]. The prefixes homo-, meso-, hetero- can be combined with the prefixes tri-, di-, or used alone (Āuroviĉ, 1994).

A tetrahedral sheet (*Tet*) can be combined with an octahedral sheet (*Oc*) either by a shared plane of apical O atoms (in all groups of hydrous phyllosilicates, Fig. 9.2.2.7a), or by hydrogen bonds (in the serpentine-kaolin group and in the chlorite group, Fig. 9.2.2.7b). Two tetrahedral sheets can either form a pair anchored by interlayer cations (in the mica group, Fig. 9.2.2.8a) or an unanchored pair (in the talc-pyrophyllite group, Fig. 9.2.2.8b).

The ambiguity in the stacking occurs at the centres between adjacent *Tet* and *Oc* and between adjacent *Tet* in the talc-pyrophyllite group. From the solved and refined crystal structures it follows that the displacement of (the origin of) one sheet relative to (the origin of) the adjacent one can only be one (or simultaneously three – for homo-octahedral sheets) of the nine vectors shown in Fig. 9.2.2.9.

The number of possible positions of one sheet relative to the adjacent one can be determined by the corresponding *NFZ* relations (§9.2.2.2.1). As an example, the contact (*Tet*; *Oc*) by shared apical O atoms, and the contacts (*Oc*; *Tet*) by hydrogen

\* A hexagonally centred unit mesh  $\mathbf{a}_1, \mathbf{a}_2$  instead of a primitive mesh  $\mathbf{A}_1, \mathbf{A}_2$  is used (Fig. 9.2.2.6b).

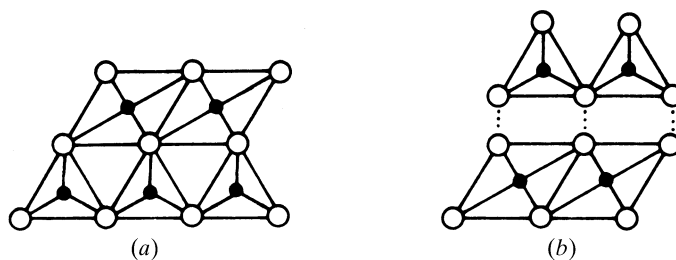


Fig. 9.2.2.7. Two possible combinations of one tetrahedral and one octahedral sheet (a) by shared apical O atoms, (b) by hydrogen bonds (side projection).

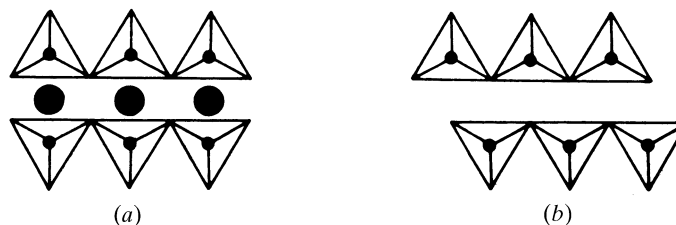


Fig. 9.2.2.8. Combination of two adjacent tetrahedral sheets (a) in the mica group, (b) in the talc-pyrophyllite group (side projection).



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bonds, for a homo-octahedral case, are illustrated in Figs. 9.2.2.10(a) and 9.2.2.10(b),(c), respectively. The two kinds of sheets are represented by the corresponding symbolic figures indicated in Fig. 9.2.2.6. For Fig. 9.2.2.10(a): the symmetry of *Tet* is  $P(3)1m$ , thus  $N = 6$ ; the symmetry of *Oc* is  $H\bar{3}12/m$  and its position relative to *Tet* is such that the symmetry of the pair is  $P(3)1m$ , thus  $F = 6$  and  $Z = 1$ : this stacking is unambiguous.\* But, if the sequence of these two sheets is reversed,  $Z = 3$ , because  $N_{Oc} = 18$  ( $h$  centring of *Oc*). For Figs. 9.2.2.10(b) and (c),  $Z = 3$ . Similar relations can be derived for meso- and hetero-octahedral sheets as well as for the pair (*Tet*; *Tet*) in the talc-pyrophyllite group.

A detailed geometrical analysis shows that the possible positions are always related by vectors  $\pm \mathbf{b}/3$ . This, together with the trigonal symmetry of the individual sheets, leads to the fact that any superposition structure (§9.2.2.5) is trigonal (also rhombohedral) or hexagonal, and the set of diffractions with  $k_{\text{ort}} \equiv 0 \pmod{3}$  has this symmetry too. This is important for the analysis of diffraction patterns.

Some characteristic features of basic types of hydrous phyllosilicates are as follows:

*The serpentine-kaolin group:* The general structural principle is shown in Fig. 9.2.2.11. The structures belong to category II (§9.2.2.7.2). In the homo-octahedral family, there are 12 non-equivalent (16 non-congruent) MDO polytypes (any two polytypes belonging to an enantiomorphous pair are equivalent but not congruent); in the meso-octahedral family, there are 36 non-equivalent (52 non-congruent) MDO polytypes. These sets are identical with the sets of *standard* or *regular* polytypes derived by Bailey (for references see Bailey, 1980) (trioctahedral) and by Zvyagin (1967) (dioctahedral and trioctahedral). The individual polytypes can be ranged into four groups (subfamilies, which are individual OD groupoid families), each with a characteristic superposition structure.

\*If the symmetry of *Tet* is  $P(6)mm$  (in the Pauling model),  $Z = 2$ . This case is common in the literature. However, with the trigonal symmetry of *Tet*, these two possibilities would correspond to Franzini (1969) types *A* and *B*, which are not geometrically equivalent.

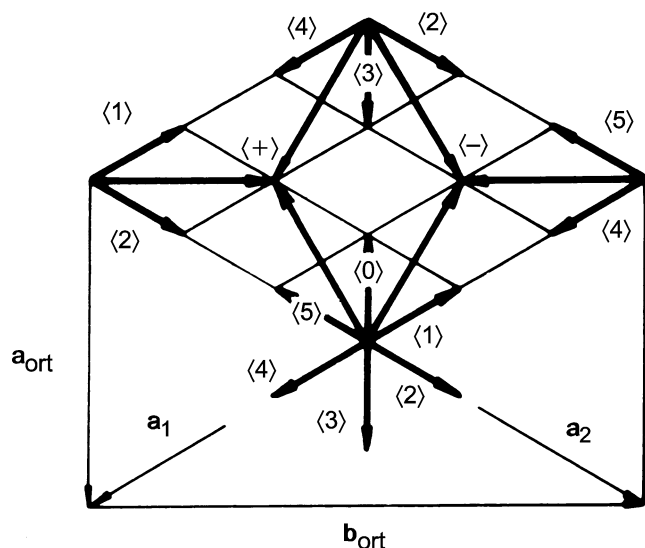


Fig. 9.2.2.9. The nine possible displacements in the structures of polytypes of phyllosilicates. The individual vectors are designated by their conventional numerical characters and the signs +, -. The zero displacement (\*) is not indicated. The relations of these vectors to the basis vectors  $\mathbf{a}_1$ ,  $\mathbf{a}_2$  or  $\mathbf{a}$ ,  $\mathbf{b}$  are evident.

*The mica group:* The general structural principle is shown in Fig. 9.2.2.12. The structures belong to category IV. There are 6 non-equivalent (8 non-congruent) homo-octahedral MDO polytypes, 14 (22) meso-octahedral, and 36 (60) hetero-octahedral MDO polytypes. The homo-octahedral MDO polytypes are identical with those derived by Smith & Yoder (1956); meso-octahedral MDO polytypes include also those with non-centrosymmetric 2:1 layers (*Tet*; *Oc*; *Tet*); some of these have also been derived by Zvyagin *et al.* (1979). The individual polytypes can be ranged into two groups (subfamilies). For complex polytypes and growth mechanisms, see Baronnet (1975, 1986).

*The talc-pyrophyllite group:* The general structural principle is shown in Fig. 9.2.2.13. The structures belong to category I. There are 10 (12) MDO polytypes in the talc family (homo-octahedral) and 22 (30) MDO polytypes in the pyrophyllite family (meso-octahedral); some of these have been derived also by Zvyagin *et al.* (1979). The structures can be ranged into two

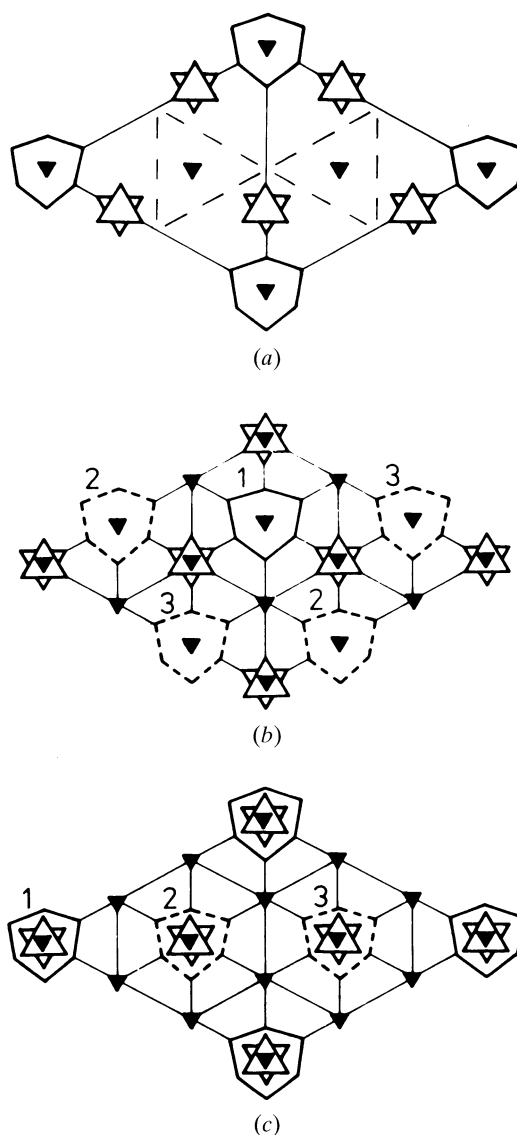


Fig. 9.2.2.10. The NFZ relations (a) for the pair tetrahedral sheet-homo-octahedral sheet (with shared apical O atoms), (b), (c) for the pair homo-octahedral sheet-tetrahedral sheet (by hydrogen bonds). The sheets are represented by their symbolic figures; some relevant symmetry elements are also indicated. One of the possible positions (labelled 1) is drawn by full, the other two (2, 3) by broken lines.

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groups (subfamilies). For more details, see also Evans & Guggenheim (1988).

*The chlorite-vermiculite group:* There are two kinds of octahedral sheets in these structures: the *Oc* sandwiched between two *Tet* and the *interlayer* (Fig. 9.2.2.14). The structures belong to category IV. Any *Oc* can be independently homo-, meso-, or hetero-octahedral, and thus, theoretically, there are nine families here. Although vermiculites have a crystal chemistry different from chlorites, they can be, from the symmetry point of view, treated together. There are 20 (24) homo-homo-octahedral, 44 (60) homo-meso-octahedral and 164 (256) meso-meso-octahedral MDO polytypes (the first prefix refers to the 2:1 layer, the second to the interlayer); the other families have not yet been treated. Some of these polytypes have also been derived by other authors (for references, see Bailey, 1980; Zvyagin *et al.*, 1979).

In order to preserve a unitary system, some monoclinic polytypes necessitate a 'third' setting, with the *a* axis unique. These should not be transformed into the standard second setting.

### 9.2.2.3.1.2. Diffraction pattern and identification of individual polytypes

Owing to the trigonal symmetry of the basic structural units and their stacking mode, the single-crystal diffraction pattern of hydrous phyllosilicates has a hexagonal *geometry* and it can be referred to hexagonal or orthohexagonal reciprocal vectors  $\mathbf{a}_1^*$ ,  $\mathbf{a}_2^*$  or  $\mathbf{a}^*$ ,  $\mathbf{b}^*$ , respectively (Figs. 9.2.2.15 and 9.2.2.16). It contains three types of diffractions:

(1) Diffractions  $00l$  (or  $000l$ ), always sharp and common to all polytypes of a family including all its subfamilies. They are indicative of the mineral group, but useless for the identification of polytypes.

(2) The remaining diffractions with  $k_{\text{ort}} \equiv 0 \pmod{3}$ , always sharp and common to all polytypes of the same subfamily.

(3) All other diffractions: sharp only for periodic polytypes, otherwise present on diffuse rods parallel to  $\mathbf{c}^*$ . These are characteristic of individual polytypes. Diffractions  $0kl$  – if sharp – are common to all polytypes of the family with the same *bc* projection.

From descriptive geometry, it is known that two orthogonal projections suffice to characterize unambiguously any structure and, therefore, the superposition structure (which implicitly contains the *ac* projection) together with the *bc* projection suffice for an unambiguous characterization of any polytype. It also follows that the diffractions with  $k \equiv 0 \pmod{3}$  together with the  $0kl$  diffractions with  $k \not\equiv 0 \pmod{3}$  suffice for its determination (except for homometric structures) (Đurovič, 1981).

From the trigonal or hexagonal symmetry of any superposition structure and from Friedel's law, it follows that the reciprocal

rows  $20l$ ,  $13l$ ,  $\bar{1}3l$ ,  $\bar{2}0l$ ,  $\bar{1}\bar{3}l$ , and  $\bar{1}\bar{3}l$  (Fig. 9.2.2.16) carry the same information. Therefore, for identification purposes, it suffices to calculate the distribution of intensities along the reciprocal rows  $20l$  (superposition structure – subfamily) and  $02l$  (*bc* projection) for all MDO polytypes. Experience shows (Weiss & Đurovič, 1980) that a mere visual comparison of calculated and observed intensities along these two rows suffices for an unambiguous identification of a MDO polytype. A similar scheme has been presented by Bailey (1988b).

The above considerations are based on the ideal Radoslovich model. Diffraction patterns of real structures may exhibit deviations owing to the distortion of the ideal lattice geometry and/or symmetry of the structure.

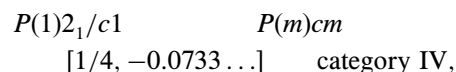
### 9.2.2.3.2. Stibivanite $\text{Sb}_2\text{VO}_5$

The crystal structure of this mineral has been determined by Szymański (1980). It turned out to be identical with that of the compound of the same composition synthesized earlier (Darriet, Bovin & Galy, 1976). The structure is monoclinic with space group  $C12/c1$ , lattice parameters  $a = 17.989$  (6),  $b = 4.7924$  (7),  $c = 5.500$  (2) Å,  $\beta = 95.13$  (3)°.

Structural units are formed of  $\text{SbO}_2\text{--O--VO--O--SbO}_2$  extended along *a*, with adjacent units bonded along *c* through  $\text{Sb--O--Sb}$  and  $\text{V--O--V}$  bonds. Ribbons are thus formed with no bonding along *b*, and only the  $\text{Sb--O}$  interactions [2.561 (4) Å] along *a* (Fig. 9.2.2.17). This accounts for the excellent acicular cleavage.

Merlino *et al.* (1989) recognized in this structure sheets of  $\text{VO}_5$  square pyramids (*Pyr*) parallel to *bc*, with layer symmetry  $P(2/m)2/c2_1/m$  alternating with sheets containing chains of distorted  $\text{SbO}_3$  tetrahedron-like pyramids (*Tet*) with layer symmetry  $P(1)2_1/c1$  (Fig. 9.2.2.18). Owing to the higher symmetry of *Pyr*, they concluded that there may also exist an alternative attachment of *Tet* to *Pyr*, such that the triples (*Tet; Pyr; Tet*) will exhibit the layer symmetry  $Pmc2_1$ , and they will be arranged so that another polytype  $2O$  with symmetry  $P2_1/m2_1/c2_1/n$  (Fig. 9.2.2.19) is formed [in the original  $2M$  polytype, the triples (*Tet; Pyr; Tet*) have the layer symmetry  $P(1)2/c1$ ]. A mineral with such a structure, with lattice parameters  $a = 17.916$  (3),  $b = 4.700$  (1),  $c = 5.509$  (1) Å, has actually been found.

The polytypism of stibivanite is reflected in its OD character: the two kinds of sheets *Pyr* and *Tet* correspond to two kinds of non-polar layers: their relative position is given by the family symbol:



the *NFZ* relations being

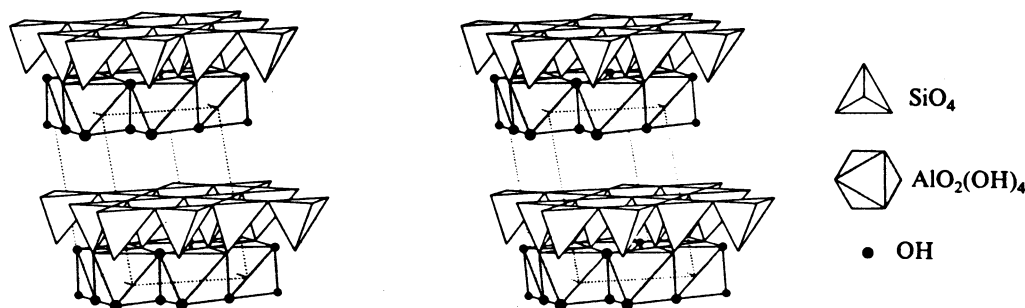


Fig. 9.2.2.11. Stereopair showing the sequence of sheets in the structures of the serpentine-kaolin group (kaolinite-1A, courtesy Zoltai & Stout, 1985).

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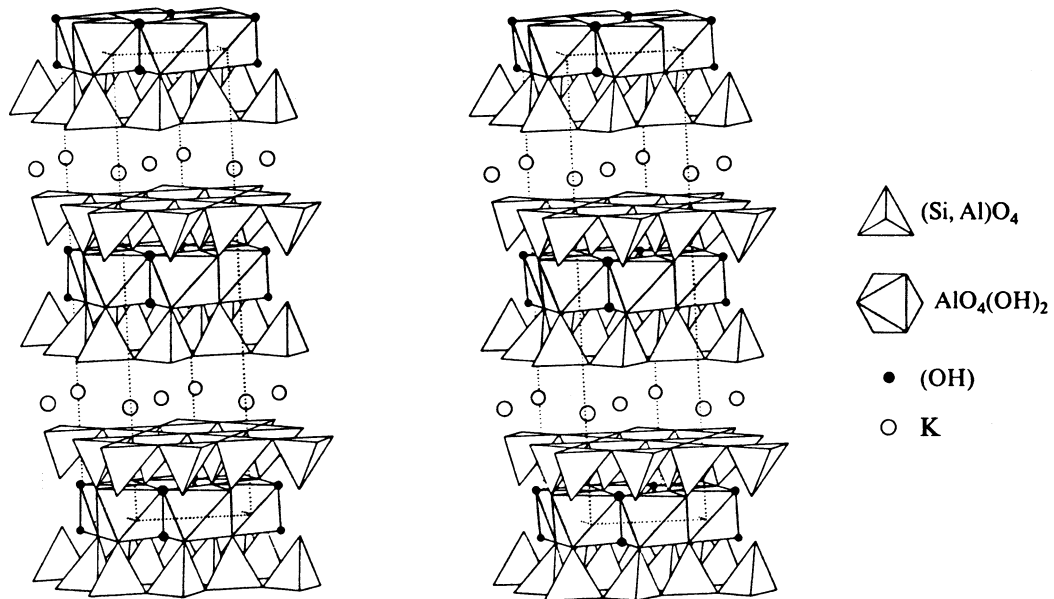


Fig. 9.2.2.12. Stereopair showing the sequence of sheets in the structures of the mica group (muscovite- $2M_1$ , courtesy of Zoltai & Stout, 1985).

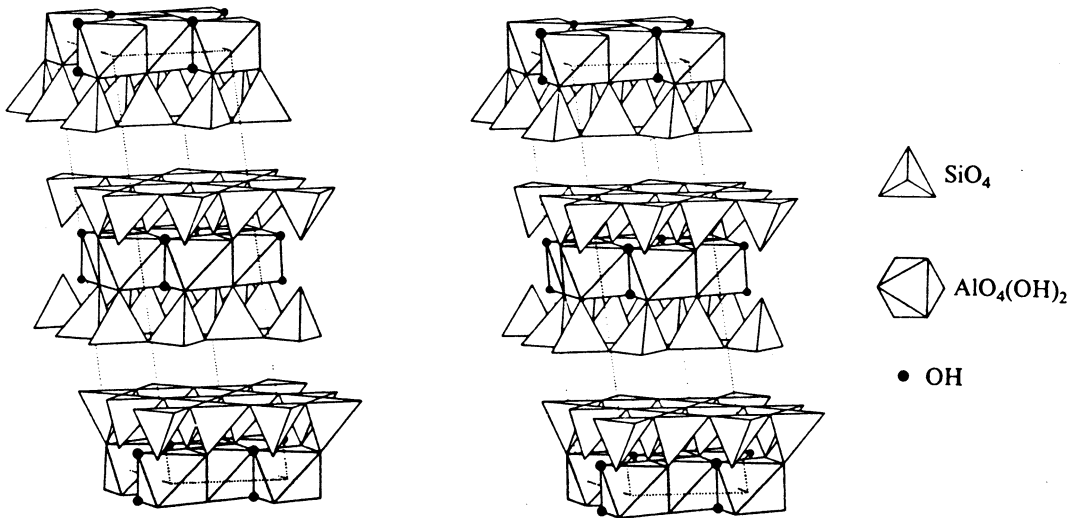


Fig. 9.2.2.13. Stereopair showing the sequence of sheets in the structures of the talc-pyrophyllite group (pyrophyllite- $2M$ , courtesy of Zoltai & Stout, 1985).

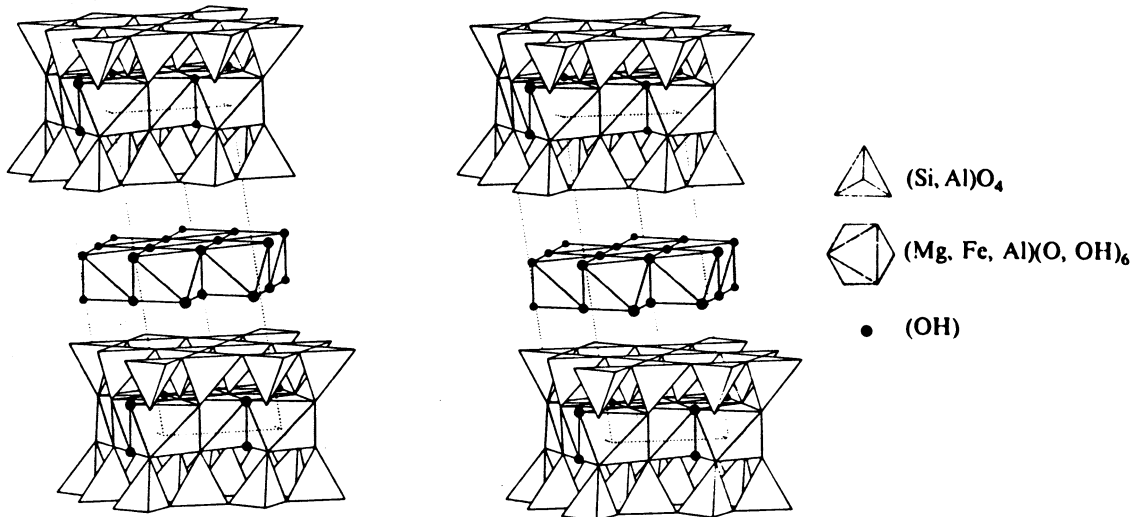


Fig. 9.2.2.14. Stereopair showing the sequence of sheets in the structures of the chlorite-vermiculite group (chlorite- $1M$ , courtesy of Zoltai & Stout, 1985).

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OD layer	layer group	subgroup of $\lambda$ - $\tau$ oper.	$N$	$F$	$Z$
$L_{2n+1}$	$P(2/m)2/c2_1/m$	$P(2)cm$	4		1
			↘	↗	
				2	
$L_{2n}$	$P(1)2_1/c1$	$P(1)c1$	2		2.
			↗	↘	

Both polytypes are slightly desymmetrized. Since the shift between the origins of *Pyr* and *Tet* is irrational, stibivanite has no superposition structure and thus the diffraction patterns of its polytypes have no common set of family diffractions. Another remarkable point is that the recognition of these structures as OD structures of layers has nothing to do with their system of chemical bonds (see above).

### 9.2.2.3.3. $\gamma$ - $Hg_3S_2Cl_2$

Among about 40 investigated crystals synthesized by Carlson (1967), not one was periodic. All diffraction patterns exhibited a common set of family diffractions, but the distribution of intensities along diffuse streaks varied from crystal to crystal (Đurovič, 1968). The maxima on these streaks indicated in some cases a simultaneous presence of domains of three periodic polytypes – one system of diffuse maxima was always present and it was referred to a rectangular cell with  $a = 9,328$  (5),  $b = 16.28$  (1),  $c = 9.081$  (6) Å with monoclinic symmetry. All crystals were more or less twinned, sometimes simulating orthorhombic symmetry in their diffraction patterns.

The superposition structure with  $A = a$ ,  $B = b/2$ ,  $C = 2c$  and space group  $Pbmm$  was solved first. Here, a comparison of the family diffractions with the diffraction pattern of the  $\alpha$

modification (Frueh & Gray, 1968) proved decisive. It turned out that only one kind of Hg atom contributed with half weight to the superposition structure. This means that only these atoms repeat in the actual structure with periods  $b = 2B$  and  $c = 2C$ . All other atoms (in the first approximation) repeat with the periods of the superposition structure and thus do not contribute to the diffuse streaks.

The symmetry of the superposition structures is compatible with the OD groupoid family determined from systematic absences:

$$\begin{matrix} A(2) & m & m \\ \{(b_{1/2} & 2_{1/2} & 2\} \\ \{(b_{1/2} & 2_{1/2} & 2\} & \text{category III,} \end{matrix}$$

where the subscripts 1/2 indicate translational components of  $b/4$  (Dornberger-Schiff, 1964, pp. 41 ff.).

The solution of the structural principle thus necessitated only the correct location of the 'disordered' Hg atom in one of two

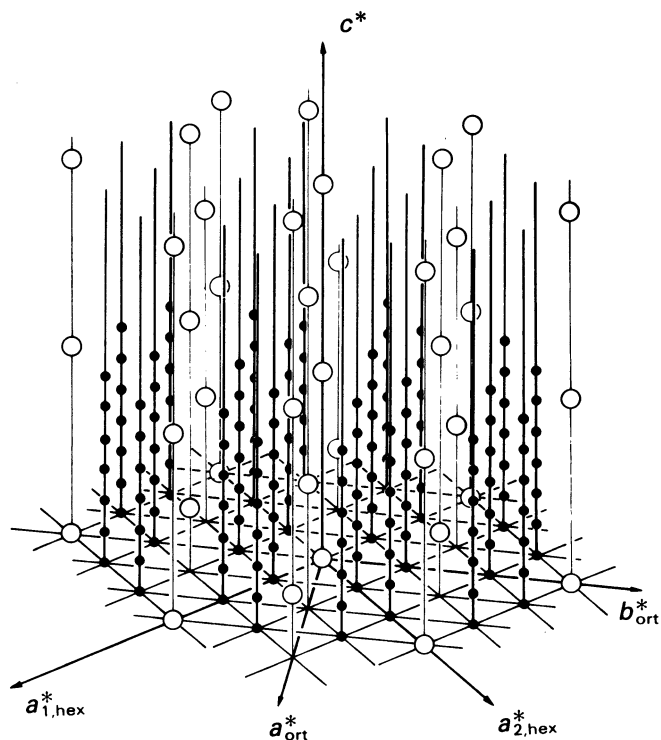


Fig. 9.2.2.15. Clinographic projection of the general scheme of a single-crystal diffraction pattern of hydrous phyllosilicates. Family diffractions are indicated by open circles and correspond in this case to a rhombohedral superposition structure. Only the part with  $l \geq 0$  is shown.

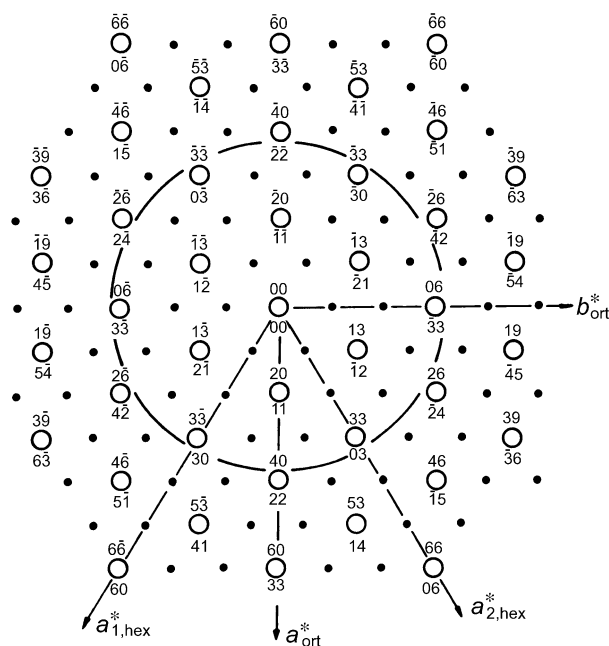


Fig. 9.2.2.16. Normal projection of the general scheme of a single-crystal diffraction pattern of hydrous phyllosilicates. Rows of family diffractions are indicated by open circles; the  $h, k$  indices refer to hexagonal (below) and orthogonal (above) axial systems.

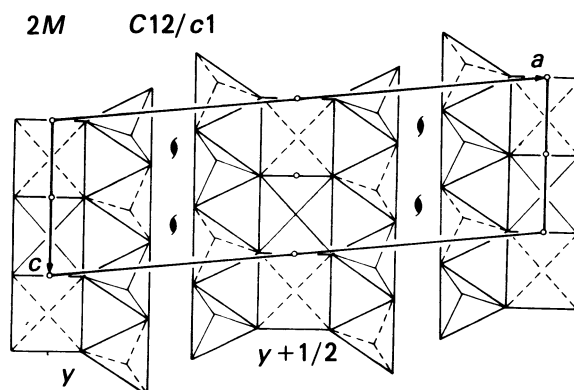


Fig. 9.2.2.17. The structure of stibivanite-2M. The unit cell is outlined and some relevant symmetry operations are indicated (after Merlino *et al.*, 1989).

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possible positions indicated by the superposition structure. An analysis of the Fourier transform relevant to the above OD groupoid family showed that the Patterson function calculated with coefficients  $|\Delta(hkl)|^2 = |F(hkl)|^2 + |F(\bar{h}kl)|^2$  shows interatomic vectors *within* any single OD layer, and it turned out that even a first generalized projection of the function yields the necessary 'yes/no' answer.

The structural principle is shown in Fig. 9.2.2.20.

There are two MDO polytypes in this family. Both are monoclinic (*c* axis unique) and consist of equivalent triples of OD layers. The first, MDO<sub>1</sub>, is periodic after two layers and has symmetry  $A2/m$ . The second, MDO<sub>2</sub>, is periodic after four layers, has symmetry  $F2/m$  and basis vectors  $2\mathbf{a}, \mathbf{b}, \mathbf{c}$ . Domains of MDO<sub>1</sub> were present in all crystals (the system of diffuse maxima, mentioned above), but some of them also contained small proportions of MDO<sub>2</sub>.

Later, single crystals of pure periodic MDO<sub>1</sub> were also found in specimens prepared hydrothermally (Rabenau, private communication). The results of a structure analysis confirmed the previous results but indicated desymmetrization (Đurovič,

1979). The atomic coordinates deviated significantly from the ideal OD model; the monoclinic angle became  $90.5^\circ$ . Even the family diffractions, which in the disordered crystals exhibited an orthorhombic symmetry, deviated significantly from it in their positions and intensities.

### 9.2.2.3.4. Remarks for authors

When encountering a polytypic substance showing disorder, many investigators try to find in their specimens a periodic single crystal suitable for a structure analysis by current methods. So far, no objections. But a common failing is that they often neglect to publish a detailed account of the disorder phenomena observed on their diffraction photographs: presence or absence of diffuse streaks, their position in reciprocal space, positions of diffuse maxima, suspicious twinning, non-space-group absences, higher symmetry of certain subsets of diffractions, *etc.* These data should always be published, even if the author does not interpret them: otherwise they will inevitably be lost. Similarly, preliminary diffraction photographs (even of poor quality) showing these phenomena and the crystals themselves should be preserved – they may be useful for someone in the future.

### 9.2.2.4. List of some polytypic structures

A few examples of some less-common polytypic structures published up to 1994 are listed below:

**Minerals:** McGillite (Iijima, 1982), tridymite (Wennemer & Thompson, 1984), pyrosmalite (Takéuchi, Ozawa & Takahata, 1983), zirconolite (White, Segall, Hutchison & Barry, 1984), Ti-biotite (Zhukhlistov, Zvyagin & Pavlishin, 1990), diamond (Phelps, Howard & Smith, 1993), scholzite (Taxer, 1992), fiedlerite (Merlino, Pasero & Perchiazzi, 1994), penkvliskite (Merlino, Pasero, Artioli & Khomyakov, 1994), lengebachite – non-commensurate structure (Makovický, Leonardsen & Moelo, 1994).

**Inorganic compounds:** Borates with general formula  $RAI_3(BO_3)_4$ , where  $R = Y, Nd, Gd$  (Belokoneva & Timchenko,

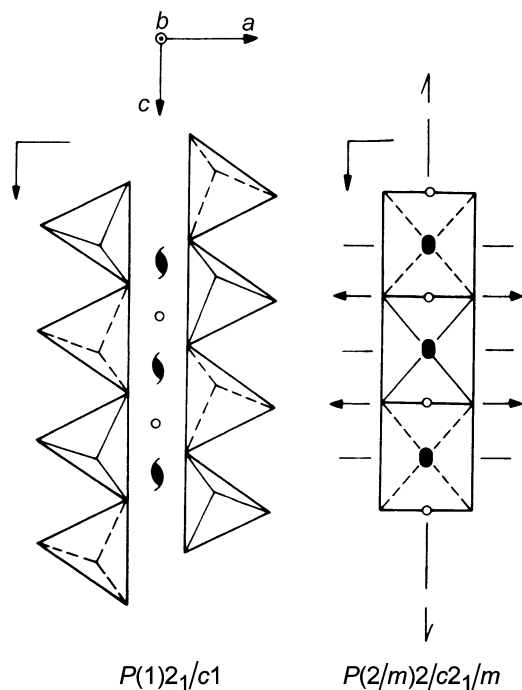


Fig. 9.2.2.18. The two kinds of OD layers in the stibivanite family (after Merlino *et al.*, 1989).

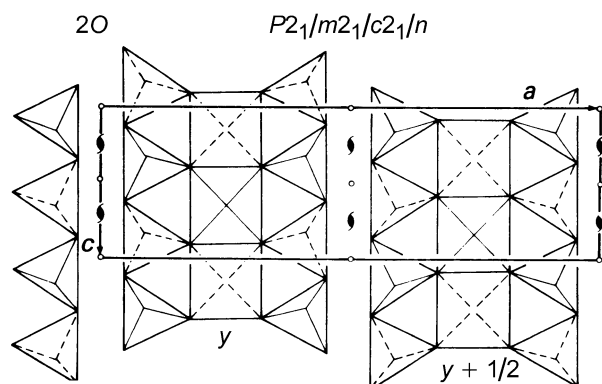


Fig. 9.2.2.19. The structure of stibivanite-2O (after Merlino *et al.*, 1989).

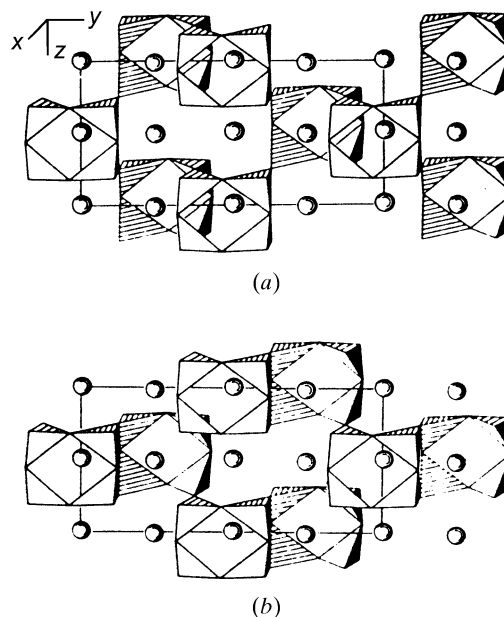


Fig. 9.2.2.20. The structural principle of  $\gamma\text{-Hg}_3\text{S}_2\text{Cl}_2$ . The shared corners of the pyramids are occupied by the Hg atoms; unshared corners are occupied by the S atoms. A pair of layers, but only the Cl atoms at their common boundary, are drawn. The two geometrically equivalent arrangements (a) and (b) are shown.

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1983), BaCrO<sub>3</sub> (Chamberland, 1983), hexacyanocomplexes of transition metals (Jagner, 1985), chromium iron carbides (Kowalski, 1985), Fe<sub>1-x</sub>S (Kuban, 1985), hexagonal copper(I) ferrite (Effenberger, 1991), PbS · 18TiS<sub>2</sub> – a modulated structure (van Smaalen & de Boer, 1992), α-LiNH<sub>4</sub>SO<sub>4</sub> (Tomaszewski, 1992), fullerene C<sub>60</sub> (de Boer, van Smaalen, Petříček, Dušek, Verheijen & Meijer, 1994).

*Organic compounds:* Oxalates (Fichtner-Schmittler, 1979), primetine (Jarchow & Schmale, 1985), 2-hydroxy-4-methoxy-2*H*-1,4-benzoxazin-3-one, C<sub>9</sub>H<sub>9</sub>NO<sub>4</sub> (Kutschabsky, Kretschmer, Schrauber, Dathe & Schneider, 1986), carbamazepine dihydrate (Reck & Dietz, 1986), piroxicam (Reck, Dietz, Laban, Günther, Bannier & Höhne, 1988), *E*-octadecanoic acid (Kaneko, Sakashita, Kobayashi, Kitagawa, Matsuura & Suzuki, 1994).

*Metals, intermetallic compounds and alloys:* Alloys are treated in a monograph by Nikolin (1984); a lecture note by Amelinckx (1986) gives details of high-resolution electron microscopy and

examples of the investigation of some alloys but also of other polytypic structures. Special papers: Li metal (Schwarz & Blaschko, 1990), Zr(FeCr)<sub>2</sub> Laves phases (Burany & Northwood, 1991), doped Co–W and Co–Mo alloys (Nikolin, Babkevich, Izdkovskaya & Petrova, 1993).

Further data are given in articles by Bailey *et al.* (1977), Dornberger-Schiff (1979), Zvyagin (1988), Baronnet (1992) and Zorkii & Nesterova (1993).

*High-resolution electron microscopy (HREM)* applied in the structure analysis of (also disordered) polytypic substances: orientite (Mellini, Merlino & Pasero, 1986), ardennite (Pasero & Reinecke, 1991), pseudowollastonite (Ingrin, 1993), laurionite and paralaurionite (Merlino, Pasero & Perchiazzi, 1993), perovskites in the system La<sub>4</sub>Ti<sub>3</sub>O<sub>12</sub>–LaTiO<sub>3</sub> (Bontchev, Darriet, Darriet, Weill, Van Tendeloo & Amelinckx, 1993), baumhauerite (Pring & Graeser, 1994), bementite (Heinrich, Eggleton & Guggenheim, 1994), parsettensite (Eggleton & Guggenheim, 1994).